

RESEARCH

Open Access



The 1831 eruption of Babuyan Claro that never happened: has the source of one of the largest volcanic climate forcing events of the nineteenth century been misattributed?

Christopher S. Garrison , Christopher R. J. Kilburn and Stephen J. Edwards

Abstract

The 1831 eruption of Babuyan Claro in the Philippines is regarded as one of the most significant volcanic climate forcing events of the nineteenth century. Modern databases have assigned the eruption a VEI of 4? and Magnitude of 4.7. Our analysis of historical sources, however, suggests that there was no such eruption in 1831 and that this date is the result of a misinterpretation of a traveller's account which had been taken to be the primary source. We therefore suggest that the 1831 eruption is a false event. In this case, one or more eruptions elsewhere must have been responsible for producing the climate-impacting stratospheric sulphate aerosol in 1831. Our results reveal the need to re-evaluate the hazard assessment of Babuyan Claro volcano and also, potentially, the quantitative treatment of the 1831 stratospheric sulphate aerosol in climate models. The Babuyan Claro example discussed in this paper therefore reinforces a call for the careful analysis of primary historical sources in volcanology.

Keywords: False eruption, Misattributed eruption, Historical, Volcanology, Climatology

Introduction

Volcanic eruptions that produce sulphate aerosols in the stratosphere are important climate forcing events (Robock 2000). A body of evidence points to such an eruption in 1831. Unusual atmospheric optical phenomena were observed around the world in August, September and October 1831, including a high haze, a dimmed blue or green sun and 'volcanic' sunsets (Arago 1832; Kiessling 1888; Symons 1888). When the Krakatoa Committee of the Royal Society of London was searching for analogous examples of the optical phenomena seen after the 1883 eruption of Krakatoa, in Indonesia, they identified those in 1831, along with those observed in 1783 after the Grímsvötn (Laki) eruption, in Iceland, as the two most similar to have been reported in the previous four centuries (Symons 1888). Anomalous surface temperatures (regionally warmer or cooler than normal) occurred

between 1831 and 1833 (White et al. 1997; Shindell et al. 2004; Fischer et al. 2007; McCarroll et al. 2013). A sulphate peak detected in ice-cores from Greenland which is ranked as either the second (Zielinski et al. 1994), third (Sigl et al. 2013) or fifth (Gao et al. 2008) largest of the nineteenth century has been dated to 1831, with deposition continuing until 1833 (Sigl et al. 2013). Even if ranked fifth (Gao et al. 2008), the only larger nineteenth century peaks are those produced by the eruptions of Tambora (1815, Indonesia), Unknown (1809), Coseguina (1835, Nicaragua) and Krakatoa.

The largest magnitude eruption listed for 1831 in the standard catalogue of historical volcanism (the Smithsonian Global Volcanism Program (GVP) database) is that of Babuyan Claro, in the Philippines, with a Volcanic Explosivity Index (VEI) of '4?' (GVP 2013). The question mark indicates a VEI value that was "particularly difficult to assign [or was] based on purely circumstantial evidence" (Simkin et al. 1981). A VEI ≥ 4 is associated with a 'definite' stratospheric injection of eruption products (Newhall

* Correspondence: c.garrison@ucl.ac.uk
UCL Hazard Centre, Department of Earth Sciences, University College London, Gower Street, London WC1E 6BT, UK

and Self 1982). This eruption has therefore commonly been inferred to be the likely source of the 1831 stratospheric sulphate aerosol (Zielinski et al. 1994; White et al. 1997; Fischer et al. 2007; Sigl et al. 2013) and, given the magnitude of the associated Greenland sulphate peak, it is regarded as one of the most significant volcanic climate forcing events of the nineteenth century, particularly so when treated as a double eruption with the 1835 (VEI = 5) eruption of Cosegüina, producing ‘enhanced’ forcing effects over a decadal time frame (Zielinski 1995; Arfeuille et al. 2014; Longpré et al. 2014; Toohey & Sigl 2017). This eruption is also listed in the Large Magnitude Explosive Volcanic Eruptions (‘LaMEVE’) database (the threshold for inclusion in which is a VEI of at least 4), with a Magnitude (M) of 4.7 (Croweller et al. 2012).

However, the record of the 1831 eruption of Babuyan Claro in the GVP database is based solely on historical observations (GVP 2013) and Zielinski (1995) has highlighted the need for the ‘thorough study’ of this ‘less well known’ eruption. In this paper we identify and analyse the relevant historical observations to investigate if they are commensurate with the suggested VEI and Magnitude (Croweller et al. 2012; GVP 2013). Our analysis shows that, to the contrary, there is no reason to believe that Babuyan Claro did erupt in 1831 and that its record in the present-day literature is the result of mistaken assumptions. We therefore suggest that the 1831 eruption of Babuyan Claro is a false event and that one or more alternative eruptions will have to be identified as the source of the 1831 stratospheric sulphate aerosol.

Method

Historical sources (typically written records) can generally be divided into two categories. ‘Primary’ sources provide first-hand evidence of a historical event, for example contemporaneous eye-witness statements. ‘Secondary’ sources later mention, discuss or interpret those primary sources. Primary sources are therefore of paramount importance in providing the evidentiary base for the secondary sources. Since the primary sources for the 1831 eruption of Babuyan Claro were not immediately identified in the GVP database (GVP 2013), we identified them according to the following method.

An initial set of sources was selected from the references and bibliographies of the successive editions of the *Volcanoes of the World* (Simkin et al. 1981; Simkin and Siebert 1994; Siebert et al. 2010; GVP 2013). The earlier sources from which they drew their information were then identified by the multi-step process illustrated in Fig. 1. In some cases, an earlier source could be identified explicitly where cited with a complete reference. In other cases its identity could be inferred, for example, on the basis of a partial reference. The multi-step process was repeated iteratively for successively earlier sources,

tracing several routes back through the literature, until the primary source(s) for the eruption had been obtained.

A total of 39 sources were analysed for the 1831 eruption of Babuyan Claro (written in English, French, German and Italian), as well as two supplementary sources from before 1831 that mention Babuyan Claro island (written in English and Spanish). These included the standard catalogues of historical volcanism by von Hoff (1841), Daubeny (1848), Landgrebe (1855), Scrope (1862), Fuchs (1865), von Humboldt (1869), Mercalli (1907), Schneider (1911), Sapper (1917, 1927), *Catalogue of the Active Volcanoes of the World* (CAVW) (1951–1975) [Part II: Philippine Islands and Cochin China (van Padang 1953)] and Macdonald (1972). Table 1 contains an entry for each source including (i) textual extracts along with translations into English where necessary; (ii) a classification as to primary (‘P’) or secondary (‘S’) source type where appropriate; and (iii) in the case of secondary sources, a list of the earlier sources (if any could be identified) on which they were based.

Results and discussion

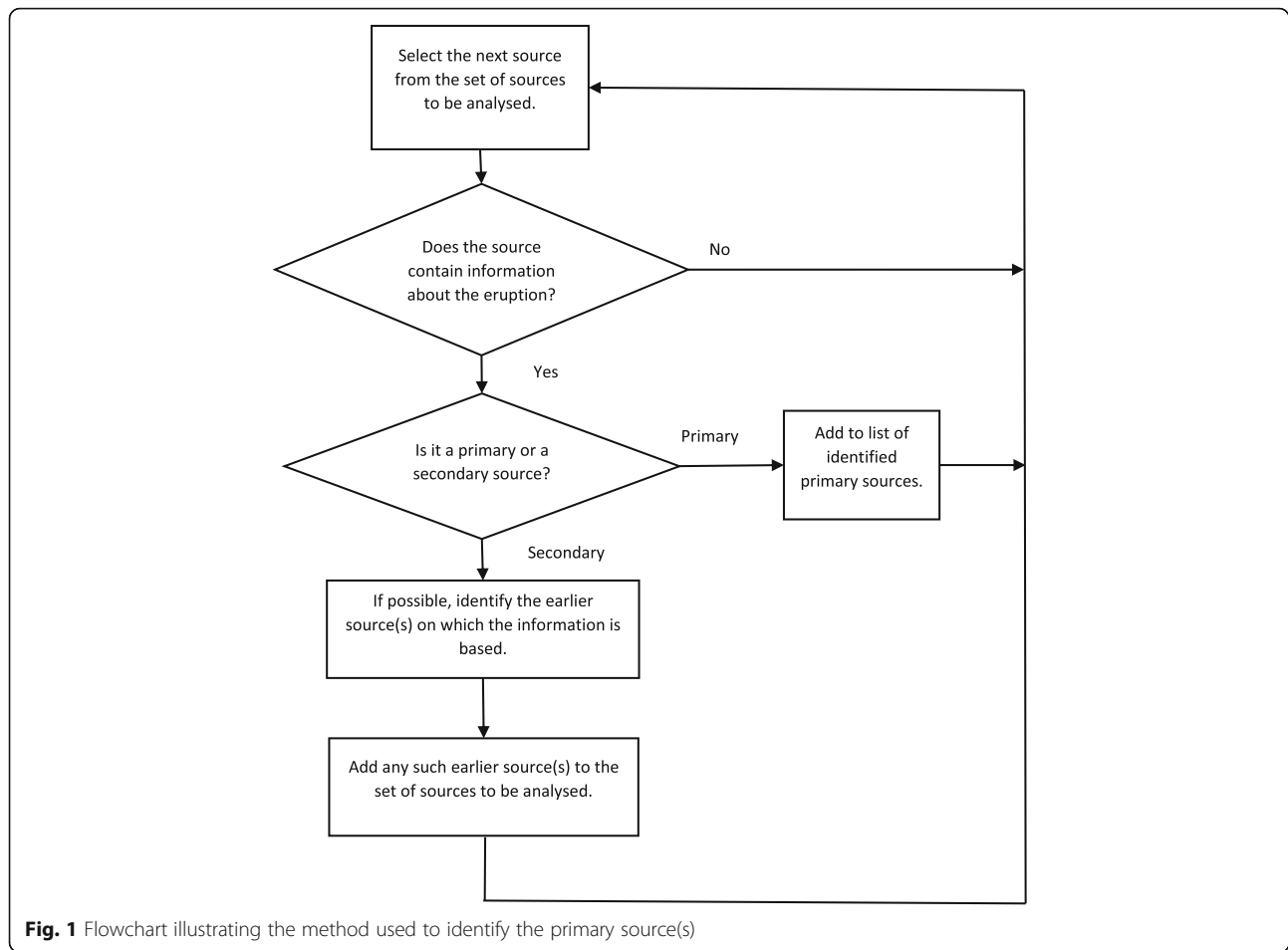
The single primary source for the 1831 eruption of Babuyan Claro

Figure 2 shows the network of routes traced back through the sources for the 1831 eruption of Babuyan Claro. A striking feature is that they all lead to Meyen (1835) as the primary source.

In 1835, Franz Julius Ferdinand Meyen (b. 1804, d. 1840) published ‘Reise Um Die Erde’ (Voyage Round the World), an account of his round-the-world voyage (1830–1832) aboard the *Prinzess Louise* (Meyen 1835; CDSB 2008). He had previously trained as a surgeon and published an important work in the field of botany (Meyen 1830; CDSB 2008).

In July 1831 the *Prinzess Louise* left the Hawaiian islands for China. The ship passed the remote island of Babuyan Claro whilst traversing the channel between the Philippines and Taiwan (Fig. 3). The island forms a part of the Luzon volcanic arc (Defant et al. 1990) and has two active volcanoes: the stratovolcano Babuyan Claro, which rises to 1080 m. a.s.l. at the centre of the island, and Smith volcano, a smaller cone at the western end of the island with a peak at 688 m. a.s.l. (GVP 2013) (Fig. 4).

On page 181 of volume 2 of his account, Meyen records that: “...We chose to approach through the Balingtang Straits and, on the following morning [7th August 1831], had the Island of Babuyan in view on the left and the Island of Balingtang on the right. *The mountains of Babuyan could exceed a few thousand feet in height, their western peak forming a steep cone, which is probably the volcano which only a short time ago had caused the inhabitants of the Island to flee.* Balingtang



island has some strangely shaped rocks in its vicinity...” (Meyen 1835, italics added; Table 1).

The western peak is Smith volcano. The italicised sentence is Meyen’s only reference to a volcano on Babuyan Claro island in his whole account. Assuming a straight line between the ship positions that Meyen noted in his meteorological log for 12 h on the 6th August and 18 h on the 7th August (Meyen 1835), Fig. 3 shows the re-constructed course of the Prinzess Louise. The reconstruction suggests that Meyen would still have been about 14 km distant at his closest approach.

Meyen evidently believed (a) that a volcanic event had taken place on the island, ‘probably’ at Smith volcano; (b) that it had caused the inhabitants of the island to flee; and (c) that it had occurred ‘only a short time ago’. However, his use of the term ‘probably’ indicates that he could not have observed any unambiguous evidence of a recent eruption. Prior to his passage past the island, his account contains no record of any of the phenomena, such as cannon-like booms, smoke columns or ash fall, which might have evidenced an eruption nearby (although, coincidentally, it appears that he did observe at least one of the unusual atmospheric optical phenomena which

were being seen around the world at the same time, see Meyen 1835; Table 1). His account includes no record of having met any of the fleeing inhabitants or, indeed, of having met anyone else who had done so. His use of the phrase ‘only a short time ago’ is difficult to interpret because elsewhere in his account he uses similar phrases to indicate intervals from hours to decades, depending on the context (Meyen 1835). Meyen’s (1835) account therefore provides no suggestion that he was reporting a volcanic event which he had witnessed himself. Rather, it is more likely that he was reporting one which he had read about (or otherwise become aware of) from an unacknowledged source describing an earlier event which had driven the inhabitants of the island to flee.

We have identified evidence pointing to at least one such event. The eighteenth century historian Vicente de Salazar reported that, in 1681, a volcano on top of a high mountain on Babuyan Claro island threw out ‘fire, rocks and ash’ and produced a ‘huge boom’ (de Salazar 1742; Table 1). Although this description points to a comparatively modest eruption, de Salazar (1742; Table 1) further reported that Fr. Matheo González, a Dominican priest visiting the island at the time, drew a vivid comparison

Table 1 Historical sources analysed for the 1831 eruption of Babuyan Claro. The sources are listed in reverse year order. All translations are one of the author's (CG) own except for the translation of source no. 41

| No. | Source | Text | Source Type: Primary / Secondary | Earlier source(s) identified. |
|-----|--|--|----------------------------------|---|
| 1 | Global Volcanism Program (2013) [Volcanoes of the World, 4th ed.] | "Start date: 1831, Stop date: Unknown, Eruption Certainty: Confirmed, VEI: 4?, Evidence: Historical Observations, Activity Area or Unit: Babuyan Claro" (Holocene Spreadsheet' downloadable from the GVP database) N.B. When viewing the eruptive history of Babuyan Claro volcano through the 'Volcano Search' option of the GVP database, a VEI = 4 (without the question mark modifier) is presented. However this is seemingly an artefact of the web interface and the VEI = 4? presented in the authoritative Holocene spreadsheet data downloadable from the GVP database is the correct value (pers. comm. Dr. Ben Andrews, Smithsonian Institution, 28/04/2016). | S | 'References' under 'General Information' |
| 2 | Siebert et al. (2010) [Volcanoes of the World, 3rd ed.] | "Start: 1831, Duration:?, Volcano Name: Babuyan Claro (Luzon-N of), Number: 0704-03=, VEI: 4?, Vol V/T: -/8" (p. 260) | S | Standard catalogues of historical volcanism (p. 1), 'References' (pp. 479 et seq) |
| 3 | PHIVOLCS (2008) | "Name: Babuyan Claro, Year: 1831, Site: Crater, Eruption Character: Explosive." | S | Van Padang (1953) |
| 4 | Simkin and Siebert (1994) [Volcanoes of the World, 2nd ed.] | "Start: 1831, Duration:?, Volcano Name: Babuyan Claro (Luzon-N of), Number: 0704-03=, VEI: 4?, Vol V/T: -/8" (p. 207) | S | Standard catalogues of historical volcanism (p. 1), 'References' (pp. 303 et seq) |
| 5 | Simkin et al. (1981) [Volcanoes of the World, 1st ed.] | "Volcano Name: BABUYAN CLARO (LUZON IS-N OF), Number: 0704-03" (p. 64) "Volcano Name: SMITH VOLCANO (LUZON IS-N OF), Number: 0704-04=, START: 1831, ERUPTIVE CHARACTERISTICS [Three symbols: Central crater eruption, normal explosions, destruction of land, property], VEI: 3?, VV LT: -8 [no recorded lava volume, 0.1 km ³ tephra" (p. 64) | S | Standard catalogues of historical volcanism (p. 1), 'Bibliography' (pp. 215 et seq) |
| 6 | COMVOL (1981) | "Name: Babuyan Claro...Its first eruption was reported in 1831." | S | Various late C19 th and C20 th sources including Van Padang (1953), Saderra Masó (1924), Smith (1924), Saderra Masó (1904). |
| 7 | MacDonald (1972) | N.B. No reference to an 1831 eruption of Babuyan Claro. | n/a | n/a |
| 8 | Lamb (1970) | "vir-viii.1831...Babuyan (Babujan), Philippine Is. (H, K, S, Sh)...19°N 122°E... (Sapper's assessment) b ₂ ...d.v./E _{max} = 300." (p. 514) "...The Babuyan eruption is generally accepted as a great one." (p. 515) | S | Humphreys (1940) Shaw (1936) Sapper (1917, 1927) Symons (1888) |
| 9 | Van Padang (1953) [CAWW, Part II: Philippine Islands and Cochin China] | "I. Name and location...BABUYAN CLARO (7,4-3)." (p. 42) "I. Name and location...SMITH VOLCANO (7,4-4)...III. Volcanic activity. Some of the older eruptions may have been of Babuyan Claro (7,4-3)...1831 [Three symbols: Eruption in the central crater, normal explosion, destruction of arable land]." (p. 43) | S | Becker (1901) Smith (1924) Saderra Masó (1905, 1924, 1925) Alvir (1928) |

Table 1 Historical sources analysed for the 1831 eruption of Babuyan Claro. The sources are listed in reverse year order. All translations are one of the author's (CG) own except for the translation of source no. 41 (*Continued*)

| No. | Source | Text | Source Type: Primary / Secondary | Earlier source(s) identified. |
|-----|------------------|---|----------------------------------|-------------------------------|
| 10 | Humphreys (1940) | "At any rate, volcanic dust is so fine, and the upper atmosphere above 11 km so free from moisture and vertical convection, that once such dust is thrown into this region, as it obviously was by the explosions of Skaptar Jökull, and Asamayama, in 1783; Babuyan, in 1831...it must require, as a rule, because of its slow descent, from 1 to 3 years to get back to the earth." (pp. 593–594) "Date: 1831–1832. ..Nature of Discrepancy: Cold...Probable Cause:...Babujan Islands, 1831" (p. 615) | S | – |
| 11 | Shaw (1936) | "Volcanic eruptions since A.D. 1800...1831 Babuyan Claro." (v.2, p. 25) | S | – |
| 12 | Alvir (1928) | "Babuyan Claro - This volcano was reported to have erupted in 1919." (p. 758) | n/a | n/a |
| 13 | Sapper (1927) | "Zum philippinischen System (1) gehören folgende Vulkanzonen: [...] 5. Babuyan Claro, 1000 m. in 19o 40' N. Br., 121o 56' ÖL., hatte 1831 heftigen A., dessen Feinaschen in Europa Dämmerungserscheinungen verursacht haben dürften (neuerdings bezweifelt). Tätig um 1860. 1917 und 1918 Anzeichen neuer Tätigkeit. [...] (1) Neue Nachrichten in M.Saderra Masó, Active Philippine Volcanoes (Bull. Weather Bureau April 1922) Manila, Auszug in Bull. volcanologique 1925, S. 306–310." (p. 320) [Translation: "The following volcanic zones belong to the Philippine system (1): [...] "5. Babuyan Claro, 1000 m....underwent a powerful eruption in 1831, whose fine ashes could have caused the twilight effects in Europe (recently questioned). [...] (1) New information in M.Saderra Masó, Active Philippine Volcanoes (Bull. Weather Bureau April 1922) Manila, Summary in Bull. volcanologique 1925, S. 306–310."] N.B. It is interesting that Sapper introduces the phrase 'recently questioned' to his 1927 entry for this eruption. He cites only Saderra Masó (1922), the summary of which presented in Saderra Masó (1925) states that a precise knowledge of the dates of historical eruptions of Babuyan Claro volcano is 'rather wanting'. However, even if Sapper did entertain any such doubts in 1927, there is no evidence to suggest that they constrained the use of his 1917 tephra volume range by later authors. | S | Saderra Masó (1922, 1925) |

Table 1 Historical sources analysed for the 1831 eruption of Babuyan Claro. The sources are listed in reverse year order. All translations are one of the author's (CG) own except for the translation of source no. 41 (*Continued*)

| No. | Source | Text | Source Type: Primary / Secondary | Earlier source(s) identified. |
|-----|--|--|----------------------------------|---|
| 14 | Saderra Masó (1925) | <p>"Claro Babuyan Volcano. On the island of Babuyan near the center at 19° 40' N, 121° 56' E. Since 1656, when Dominican Missionaries visited the island for the first time, it has been reported as active but precise dates of eruptions are rather wanting. In 1907 WORCESTER took photographs which show two new recent craters. In 1917, W. D. SMITH of the Bureau of Mines, visited it and found convincing indications of recent outbursts. More recently, May 17th and 19th, 1918, Captain ROSÉS of the interisland S. "Mauban" on sailing close to the island saw the volcano in eruption. West of the main volcano rises another remarkable well preserved cone called Smith volcano, which was active within the past six months. Several of the volcanoes appearing in previous lists, as for instances, those published by PERRY, MARCALLI and others, are extinct very old volcanoes of the worm-down-stock type..." (pp. 385–386)</p> | n/a | Mercalli (1907) Perrey (1860) |
| 15 | Saderra Masó (1924) (see Saderra Masó 1925) | | | |
| 16 | Smith (1924) | <p>"On the way, I was enabled to remain about two hours on Babuyan Claro, where there is a volcano which has been active as recently as 1860. There are two volcanoes on this island, one a beautifully symmetrical cinder cone, about 2220 ft [670 m] in height, and the other less symmetrical but showing two craters which contain steam vents. The smaller volcano contains, besides the fine ash that gives it its shape, a recent flow of scoriaceous basalt. The older flows of this volcano are also basaltic. There has been no serious eruption since 1860. The few people living on the island report, however, that steam explosions occurred on the smaller mountain ten years ago and on the larger four years ago. No ashes reached the village in any of these explosions, however. As it is only in the last thirty-five years that the island has been inhabited, the inhabitants know nothing of the eruptions of 1860." (p. 264)</p> <p>"BABUYAN CLARO. Ferguson says: Babuyan Island is about 13 km in a northeast and southwest direction and has an average width of about 10 km. At the western point is a volcano 670 m high. This mountain is conical in shape and evidently contains a small crater at its summit. Flows of rough scoriaceous basalt surround the base and the mountain itself is built up of alternating basalt flows and deposits from explosive eruptions, angular basalt fragments, bomb lapilli and small slaglike masses. The writer had an opportunity to land here for a</p> | S | Becker (1901) Fuchs (1881) [Meyen (1835)] |

Table 1 Historical sources analysed for the 1831 eruption of Babuyan Claro. The sources are listed in reverse year order. All translations are one of the author's (CG) own except for the translation of source no. 41 (*Continued*)

| No. | Source | Text | Source Type: Primary / Secondary | Earlier source(s) identified. |
|-----|--|--|----------------------------------|--|
| 17 | Saderra Masó (1922) (see Saderra Masó 1925) | short time and, although unable to reach the summit, collected specimens of the basalt from one of the earlier flows and from the latest flow. In the eastern part of the island is another volcano, 1160 m in height and heavily timbered. On its southern side were active fumaroles. Several eruptions of the volcanoes of this island have been noted. | | |
| 18 | Sapper (1917) | Horsburgh in 1817 mentions a volcano on the western point, though he does speak of any eruption having occurred. Fuchs [1881] mentions an eruption as having occurred in 1831. Becker quotes Meyen as mentioning this eruption, and Semper as stating that Babuyan Claro seemed then (1860) to be continually in eruption." (p. 304) NB. It is interesting that, consistent with Maree (2005), Smith became aware on his brief visit to Babuyan Claro island that it had been uninhabited till comparatively recently (although he seems to have been unaware that it had been previously inhabited in the more distant past). | S | Saderra Masó (1904) Symons (1888) Meyen (1835) |
| | | "Babuyan Claro, 1000 m, hatte 1831 (5) einen heftigen Ausbruch, dessen Feinaschen noch in Europa schöne Dämmerungserscheinungen mitverursacht haben dürften (6). [...] (5) Masó a) S. 12. Meyen, Reise II, S. 181. (6) Report of the Krakatoa Committee of the Royal Society. London 1888. S.396." (p. 149) [Translation: "Babuyan Claro, 1000 m, underwent a powerful eruption in 1831 (5) the fine ashes of which could have contributed to the beautiful twilight effects even in Europe (6). [...] (5) Masó a) P. 12. Meyen, Reise II, P. 181. (6) Report of the Krakatoa Committee of the Royal Society. London 1888. P.396."] the first reference is to Saderra Masó (1904). "Tabelle der bekannten Riesenausbrüche... 1831 Babuyan Claro b ₂ " (pp. 338-340) [Translation: "Table of known large eruptions... 1831 Babuyan Claro b ₂ "] "a ₁ a ₂ bedeutet darin wieder Förderung von über 1 bzw. 1/10 cbkm Lava, b ₁ b ₂ von ebensoviel Lockermassen." (p. 337) [Translation: "In the following a ₁ a ₂ again denotes | | |

Table 1 Historical sources analysed for the 1831 eruption of Babuyan Claro. The sources are listed in reverse year order. All translations are one of the author's (CG) own except for the translation of source no. 41 (*Continued*)

| No. | Source | Text | Source Type: Primary / Secondary | Earlier source(s) identified. |
|-----|---|--|----------------------------------|---|
| 19 | Schneider (1911) | the production of over 1 or 1/10 km ³ of lava, b ₁ b ₂ of the same amount of tephra.]" "90...Babuyan Claro...121°56'3. 19°30' n...Ausbrüche: 1831, 1860." (p. 242) [Translation: "90...Babuyan Claro...Eruptions: 1831, 1860"] | S | – |
| 20 | Mercalli (1907) | "Babuyane-Claro (1000 m.), situato nelle isole Babuyane, che formano la parte più settentrionale dell'arcipelago: ebbe un'eruzione molto forte nel 1831." (p. 305) [Translation: "Babuyane-Claro (1000 m.), located in the Babuyan islands, which form the northernmost part of the archipelago: underwent a powerful eruption in 1831."] | S | Perrey (1860) |
| 21 | Saderra Masó (1905) | "[Name] Babuyán Claro...[Date of eruption] 1831, 1860" (p. 187) | S | – |
| 22 | Saderra Masó (1904) | "Name: Babuyán Claro...Date of eruption: 1831, 1860" (p. 12) | S | – |
| 23 | Becker (1901) | "The remaining volcanic mountains of the archipelago lie in one group at its northern end. Mr. James Horsburgh mentions the little island Camaguin de Babuyanes as having formerly been a volcano, and says that on the west end of the island of Babuyan Claro there is a volcano [ref: Horsburgh 1817, p. 328]. Meyen states that in 1831 the latter underwent a violent eruption. Semper [see Semper (1869)] says Babuyan Claro seems to be continually in eruption..." (p. 54) | S | Semper (1869) Meyen (1835) Horsburgh (1817) |
| 24 | Symons (ed.) (1888) [Report of the Krakatoa Committee of the Royal Society] | "Babujan Islands, 1831, (Great eruption)" (p. 396) | S | Judd (1881) Daubeny (1848) |
| 25 | Fuchs (1884) | "On connaît plusieurs volcans dans les petites îles Bajuban, qui sont situées au nord de l'archipel. L'un d'entre eux, le Claro Babuyan (10° 27' lat. Bor. et 110° 42' longit. or), est toujours actif et eut une éruption en 1831." (p. 236) [Translation: "Several volcanoes are known in the Bajuban isles, which are located in the north of the archipelago. One of them, Claro Babuyan... is still active and underwent an eruption in 1831."] | S | – |
| 26 | Fuchs (1881) (see Fuchs 1865, 1884) | | | |
| 27 | Judd (1881) | N.B. No reference to an 1831 eruption of Babuyan Claro. | n/a | n/a |

Table 1 Historical sources analysed for the 1831 eruption of Babuyan Claro. The sources are listed in reverse year order. All translations are one of the author's (CG) own except for the translation of source no. 41 (*Continued*)

| No. | Source | Text | Source Type: Primary / Secondary | Earlier source(s) identified. |
|-----|--|--|----------------------------------|--|
| 28 | Von Humboldt (1869) (see von Humboldt 1858) | | | |
| 29 | Semper (1869) | N.B. No reference to an 1831 eruption of Babuyan Claro although Semper mentions "...ein wie es scheint in beständiger Eruption befindliche Vulcan auf Babuyan Claro..." (p. 14) [Translation: "...a volcano on Babuyan Claro which seems to be in constant eruption..."] | n/a | - |
| 30 | Fuchs (1865) | "Zwischen Formosa und den Philippinen liegen zwei vulkanische Inseln, Claro Babyan und Caminguin. Der Vulkan auf Claro Babyan liegt 19° 27' n. Br., 119 ° 42' östl. L. und hat nach Meyen 1831 einen Ausbruch gehabt." (p. 47) [Translation: "Between Formosa and the Philippines lie two volcanic islands, Claro Babyan and Caminguin. The volcano on Babuyan Claro... according to Meyen, underwent an eruption in 1831."] | S | Meyen (1835) |
| 31 | Scrope (1862) | "Some of the lesser islands which connect Formosa with the Philippines have been seen in eruption. In this latter group no less than nineteen lofty insulated conical mountains, all called in the country 'volcanes,' are enumerated by Von Buch." (p. 460) | n/a | n/a |
| 32 | Perrey (1860) | "1831 - Le volcan situé à la pointe méridionale de l'île Claro Babuyan (Philippines), entre les îles Bashi et Luçon, par 19° 27' lat. N. et 119 ° 42' long. E., eut une éruption si violente, que, pour échapper à l'embracement, les habitants de l'île furent forcés de s'enfuir avec la plus grande rapidité (3). Les petites îles Baschi, et les Babuyanes, qui, dit Humboldt (4), suivant le témoignage de Meyen, ont donné encore en 1831 le spectacle d'une violente éruption de flammes..." [...] "(3) Landgrebe, l.c., p. 348, d'après Meyen, Reise um die Erde, t. II, p. 184. De Buch, l.c., p. 438, même source. (4) Cosmos, t. IV, p. 421" (p. 180) [Translation: "1831 - The volcano located at the southern tip of Claro Babuyan island (Philippines), between the Bashi islands and Luzon,..., underwent an eruption so violent that, in order to escape the blaze, the inhabitants of the island were forced to flee with the greatest rapidity (3). The small Baschi islands, and the Babuyanes, which, states Humboldt (4), according to the testimony of Meyen, yielded in 1831 the sight of a violent eruption of flames..." [...] | S | Von Humboldt (1859) Landgrebe (1855) Von Buch (1836) [Meyen (1835)] |

Table 1 Historical sources analysed for the 1831 eruption of Babuyan Claro. The sources are listed in reverse year order. All translations are one of the author's (CG) own except for the translation of source no. 41 (*Continued*)

| No. | Source | Text | Source Type: Primary / Secondary | Earlier source(s) identified. |
|-----|---|---|----------------------------------|-------------------------------|
| 33 | Von Humboldt (1859) (see von Humboldt 1858) | "(3) Landgrebe, l.c., p. 348, according to Meyen, Reise um die Erde, v. II, p. 184. De Buch, l.c., p. 438, same source. (4) Cosmos, v. IV, p. 421" | | |
| 34 | Von Humboldt (1858) | "Die kleinen Baschi-Inseln und die Babuyans, welche noch 1831 nach Meyen's Zeugniß einen heftigen Feuerbruch erlitten, verbinden Formosa mit den Philippinen..." (v. 4, p. 404) [Translation: "The small Baschi-islands and the Babuyans, which according to Meyen's testimony underwent a powerful fiery eruption in 1831, connect Formosa with the Philippines..."] | S | Meyen (1835) |
| 35 | Landgrebe (1855) | "8. Die Insel Claro Babuyan. Sie liegt in der Mitte zwischen den Bashi-Inseln und Luzon, oberhalb Camiguin. Auf ihrer Südspitze, unter 19° 27' n. Br. und 119° 42' östl. L., bemerkt man einen mehrere tausend Fuss hohen Vulkan, der nach Meyen's Zeugniß (s. dessen Reise Um die Erde Bd. II. S. 184) im J. 1831 einen so heftigen Ausbruch hatte, dass die Bewohner der Insel sich zur schnellsten Flucht genöthigt sahan, um dem sichern Verderben zu entgehen." (p. 348) [Translation: "8. Babuyan Claro island. This is located in the middle between the Bashi-islands and Luzon, above Camiguin. At its southern point...a volcano several thousand feet high can be seen, which according to Meyen's testimony (see his Voyage around the World, Volume II, Page 184) underwent such a powerful eruption in 1831, that the inhabitants of the island had to flee as quickly as possible, in order to escape from their certain destruction."] N.B. The reference to v. 2, p. 184 of Meyen (1835) is evidently a mistaken reference to p. 181. It is interesting to note the degree of embellishment which has been introduced here: not only did the inhabitants have to flee the island but they had to do so 'as quickly as possible'. | S | Meyen (1835) |
| 36 | Daubeny (1848) | "The volcanic chain is also connected more closely with the island of Formosa by a burning mountain existing in the group of the Babujan islands intermediate, in which a great eruption that occurred in 1831 drove the inhabitants from the island." (p. 399) | S | Von Buch (1836) |
| 37 | Von Hoff (1841) | N.B. No reference to an 1831 eruption of Babuyan Claro (although at v.2, p. 182, von Hoff does refer | n/a | n/a |

Table 1 Historical sources analysed for the 1831 eruption of Babuyan Claro. The sources are listed in reverse year order. All translations are one of the author's (CG) own except for the translation of source no. 41 (*Continued*)

| No. | Source | Text | Source Type: Primary / Secondary | Earlier source(s) identified. |
|-----|-----------------|--|----------------------------------|-------------------------------|
| 38 | Von Buch (1836) | to Meyen's Reise um die Welt [sic]. "17° Volcan de l'île de Babujan, Lat. 19 ° 27' N.; Long. 119 ° 42' 1/4. E. de Paris. Ce volcan qui a quelques mille pieds de hauteur, est, situé dans la partie ouest de l'île. Une grande éruption qui eut lieu en 1831, força les habitants à prendre la fuite et à abandonner l'île (Meyen, Voyage, II, 181)." (p. 438) [Translation: "17 ° Volcano on the island of Babujan... This volcano, which is some several thousand feet high, is located at the western end of the island. A great eruption, which took place in 1831, forced the inhabitants to take flight and abandon the island (Meyen, Voyage, II, 181)."] | S | Meyen (1835) |
| 39 | Meyen (1835) | "In der Nacht zum 7ten August näherten wir uns den Bashee-Inseln, deren Länge noch auf allen Karten, die wir an Bord der Prinzess hatten, verschieden angegeben war. Wir wählten die Strasse Balingtang zur Durchfährt und hatten am folgenden Morgen die Insel Babuyan zur Linken, und die Insel Balingtang zur Rechten im Gesicht. Die Berge von Babuyan können einige Tausend Fuss an Höhe übersteigen, ihre westliche Spitze bildet einen schroffen Kegelberg, der wahrscheinlich der Vulcan ist, welcher noch vor kurzer Zeit die Bewohner der Insel zur Flucht getrieben hatte. Die Balingtang-Insel hat in ihre Nähe sehr merkwürdig gestaltete Klippen..." (v. 2, p. 181) [Translation: "During the night before the 7th August, we approached the Bashee-islands, the latitude of which was recorded differently on all the maps aboard the Princess. We chose to approach through the Balingtang Straits and, on the following morning, had the island of Babuyan in view on the left and the island of Balingtang on the right. The mountains of Babuyan could exceed a few thousand feet in height, their western peak forming a steep cone, which is probably the volcano which only a short time ago had caused the inhabitants of the island to flee. Balingtang island has some strangely shaped rocks in its vicinity..."] N.B. On his approach to Babuyan Claro island, Meyen reported the observation of a solar halo and an extended twilight: "Am 3ten August... Kurz vor Mittags-Zeit, als uns die Sonne fast im Zenith stand, bildete sich ein Dunstring um dieselbe, dessen Radius 21 1/2° betrug; die Farbe des Ringes war nicht so bestimmt, wie die eines Regenbogens, und im Innern der Ringes befand sich eine Wolkenmasse, welche dunkler war, als die des umgebenden Himmels, durch welche wir die Sonne kaum durchsehen konnten. Es befand sich Niemand an Bord des Schiffes, dem ein solcher Sonnen-Ring, welcher etwas Aehnlichkeit mit dem | P / S | n/a |

Table 1 Historical sources analysed for the 1831 eruption of Babuyan Claro. The sources are listed in reverse year order. All translations are one of the author's (CG) own except for the translation of source no. 41 (*Continued*)

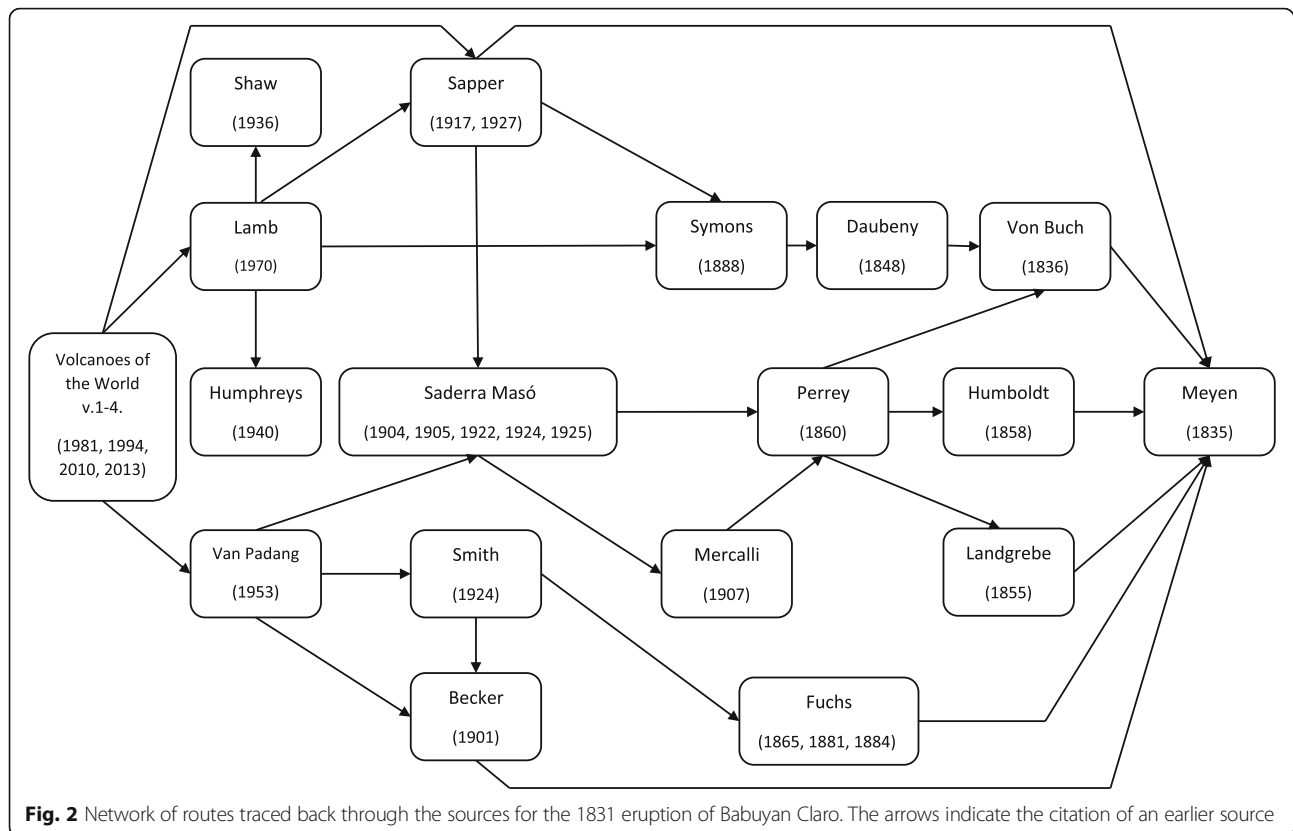
| No. | Source | Text | Source Type: Primary / Secondary | Earlier source(s) identified. |
|-----|--------|--|----------------------------------|-------------------------------|
| | | <p>sogenannten Hofe des Mondes hatte, zu Gesicht gekommen war." (v.2, pp. 180–181)</p> <p>[Translation: "On the 3rd August... Shortly before Mid-day, with the Sun almost at the Zenith for us, a hazy Ring appeared around the same, whose radius was $21 \frac{1}{2}^{\circ}$, the colour of the Ring was not so well-defined as that of a Rainbow, and inside the Ring there was located a Cloud-mass which was darker than that of the surrounding sky, through which we could hardly see the Sun. There was no-one on board the Ship who had observed such a Sun-Ring before, which had some similarity with the so-called Moon Halo."] "Auch wurde in diesen Tagen eine Dämmerung bemerkt, die beinahe eine Stunde nach Sonnen-Untergang anhielt, woran wohl die niedere Declination der Sonne, nämlich gerade 16°, den grössten Antheil hatte; diese Dämmerung wurde aber auch nur wenige Tage hindurch beobachtet." (v.2, p. 181)</p> <p>[Translation: "A Twilight was also experienced in these days which continued almost an hour after Sun-set, which was mainly caused by the low declination of the Sun, namely 16°; this Twilight was only observed for a few days however."] A solar halo known as a 'Bishop's Ring' is sometimes observed after a volcanic eruption (Meinel and Meinel 1983). Caused by the diffraction of sunlight by volcanic aerosol particles, it typically appears as a brighter whitish-blue disc surrounding the sun, enclosed by a dimmer reddish-brown ring: the inner and outer radius of the ring varies by several degrees depending on the particular size distribution of the aerosol but average values are c. $10\text{--}11^{\circ}$ and $22\text{--}23^{\circ}$ respectively (Symons 1888; Meinel and Meinel 1983). Meyen's description seems, however, a closer match to a more common 22° solar halo. Caused by the refraction of sunlight (or moonlight) by atmospheric ice-crystals, this appears as a brighter rainbow-like arc around a dimmer portion enclosing the sun, with a well defined reddish inner radius at 21.7° extending outward to a more diffuse bluish-white edge (Lynch and Livingston 2001). The circumscribed 22° solar halo which is observed at high solar elevations is particularly pronounced (Können 2015). The extended twilight appears to be an observation of the unusual atmospheric optical phenomena being contemporaneously observed elsewhere in the world at the time (as mentioned in the Introduction) but is therefore not local to Babuyan Claro island or even to East Asia (Arago 1832; Kiessling 1888; Symons 1888).</p> | | |

Table 1 Historical sources analysed for the 1831 eruption of Babuyan Claro. The sources are listed in reverse year order. All translations are one of the author's (CG) own except for the translation of source no. 41 (*Continued*)

| No. | Source | Text | Source Type: Primary / Secondary | Earlier source(s) identified. |
|-----|-------------------|--|----------------------------------|-------------------------------|
| 40 | Horsburgh (1817) | "CLARO BABUYAN, or OLD BABUYAN, in lat. 19° 37' N., lon. 122° 17' E., distant about 10 leagues to the eastward of Calayan, is the most northerly and highest of these islands, in extent about 2 or 2 1/2 leagues. There is a reef projecting from the West end of the island, and the mount on this part is a volcano; betwixt which, and the mountains on the eastern part, there is a concave curve in the form of a crescent, when viewed from the northward or southward; but when the island is seen at a great distance from the eastward, it appears as one round mountain, with a detached hummock to the northward." (p. 328) | n/a | n/a |
| 41 | De Salazar (1742) | "CAP. XXIII. FUNDACION DE LA MISSION de Battanes, y muerte de los Padres Fr. Matheo Gonzalez y Fr. Juan Rois...El año de 1680...el año siguiente...Haban sumamente afligidos, y atemorizados con el mucho suego, piedras, y cenizas, que actualmente arrojaba de si un Volcan, que ay en un monte alto de dicha Isla, que à todos causaba terrible horror, y espanto; y con la ocasion de la pena, en que se hallaban les comenzó à predicar el Padre Fr. Matheo, explicandoles las penas de el inferno, y lo incomprehensible de los tormentos, que han de padecer los condenados por una eternidad, abrassandose en sus vorazes llamas, especialmente los que siendo Christianos, avian apostatado de la Feè, y eran causa de que se condenassen sus hijos, y nietos, por tenerlos en aquella Isla, tan apartados de la vista, y comunicacion de el Ministro, que pudiesse doctrinarlos, y bautizarlos. Estando actualmente predicando el Padre Fr. Matheo, era grande el estruendo, y ruido de el Volcan, como de gruesos tiros de artilleria, con lo qual fue tan grande la conmocion, que la doctrina de el Padre causò en todos ellos, que arrassados en lagrimas sus ojos, dieron clarissimas muestras de un grande arrepentimiento de sus culpas, y se rindieron à la voluntad de el Padre Fr. Matheo, determinandos à dexar aquella Isla, y acompañar al dicho Religioso, para que los restituesse à la Iglesia, como lo executaron puntualmente, sin quedar ni uno solo en toda la Isla. Vinose el Padre con ellos à Cagayan..." (pp. 518–520) | n/a | n/a |
| | | [Translation (Montse Manzano Fernandez): "Chapter XXIII. Foundation of the Mission of Battanes, and death of the Fathers Fr. Matheo Gonzalez and Fr. Juan Rois...Year 1680...the following year... People were worried and frightened by a volcano, which was on top of a high mountain and that | | |

Table 1 Historical sources analysed for the 1831 eruption of Babuyan Claro. The sources are listed in reverse year order. All translations are one of the author's (CG) own except for the translation of source no. 41 (*Continued*)

| No. | Source | Text | Source Type: Primary / Secondary | Earlier source(s) identified. |
|-----|--------|--|----------------------------------|-------------------------------|
| | | <p>was throwing out fire, rocks and ash. This volcano caused everyone a dreadful horror. With the occasion of their sadness, Father Fr. Matheo started preaching to them, explaining the punishments of Hell, and the incomprehensible torments that those condemned for eternity would endure, burning in voracious flames. Especially those who were Christians and had abandoned (had apostatized from) their faith. For this reason they had also condemned their children and grandchildren, and for having them isolated on the island, far from the sight and the reach of a Minister who could indoctrinate and baptise them. One day, when Father Fr. Matheo was preaching, there was a huge boom and noise from the volcano, as if it was heavy artillery fire. People were so moved by the preaching from the priest that, with their eyes full of tears, they showed very clear signs of a great repentance for their sins, and did what Father Matheo wanted. They were determined to leave that Island, and to accompany the priest, so that they could be included in the Church again. They left straight away and there was no one left on the whole island. Father Fr. Matheo went with them to Cagayan....]</p> | | |



for the inhabitants between the observed phenomena and the torments of those condemned to Hell, whereupon they all decided to return with the priest to the mainland in order to remain in communion with the Church. Whether this is a wholly objective account is unclear. Based on the historical and genealogical research that she has conducted with the Ibatan people living on Babuyan Claro, Maree (2005) also suggests that the 1681 eruption led to the depopulation of the island and that it remained uninhabited thereafter until the second half of the nineteenth century. Additional research will be necessary to investigate this evidence further. Clearly, it would seem doubtful that Meyen (1835) would knowingly have considered a span of 150 years to be ‘only a short time ago’ but whether the unacknowledged source from which he gleaned his information reliably represented the date of the event described is as yet unknown.

For present purposes, it can be concluded that Meyen’s (1835) account does not represent a primary source and cannot provide any evidence for an eruption of either volcano on Babuyan Claro island in 1831.

The 1831 eruption of Babuyan Claro in the later literature

If this is the case, the question arises as to how an 1831 eruption of Babuyan Claro came to appear in the later literature and, in particular, how it came to be represented with a VEI of 4? and Magnitude of 4.7

(Croweller et al. 2012; GVP 2013). Inverting the set of routes from the present day back to Meyen (1835) (Fig. 2) suggests several possible pathways from 1831 forward to the present day. In fact, the present day representation of the 1831 eruption can plausibly be explained taking into account the steps taken by just two authors.

Leopold von Buch published an eruption catalogue in 1836, one year after the publication of Meyen (1835). He included the following entry: “Volcano on the island of Babujan [sic]...This volcano, which is some several thousand feet high, is located at the western end of the island. A great eruption, which took place in 1831, forced the inhabitants to take flight and abandon the island (Meyen, Voyage, II, 181).” (Von Buch 1836; Table 1). Meyen’s 1835 account is the only source that Von Buch cites to support this entry and yet the two are markedly different. Von Buch evidently regarded the passage as pointing to a volcanic eruption and seems to have made the following implicit assumptions: (A1) that it must have been a ‘great’ eruption, otherwise the inhabitants would not have fled; and (A2) that ‘a short time ago’ must arbitrarily mean it took place in 1831. Von Buch’s (1836) catalogue was influential and, based on this interpretation of Meyen’s account, an entry for a fictitious ‘great’ eruption on Babuyan Claro island in 1831 was widely included in the subsequent literature.

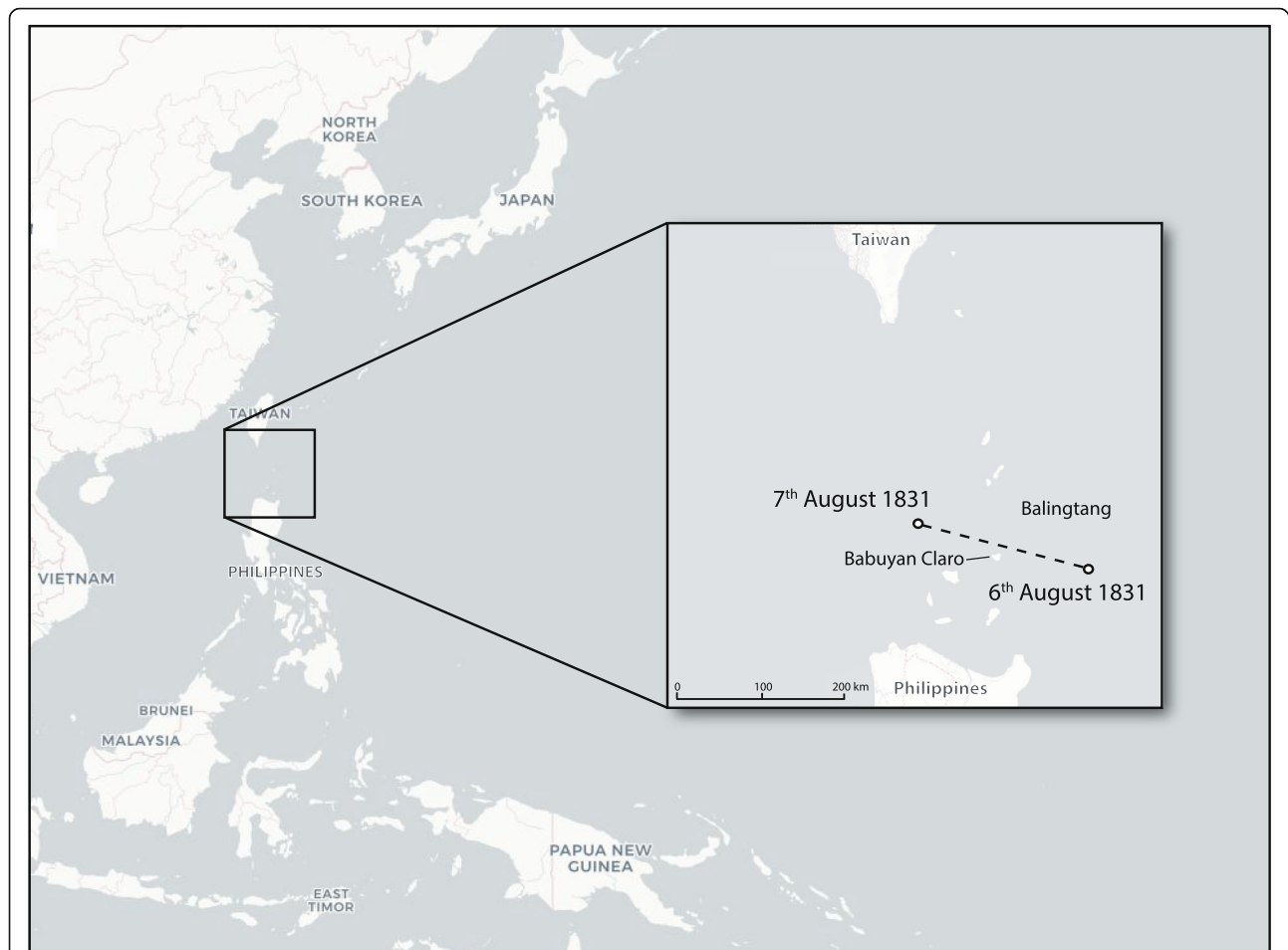


Fig. 3 The re-constructed course of the Prinzess Louise passing Babuyan Claro island in August 1831. The use of this static ArcGIS map image is permitted in academic publications (including research journals) with the following source attribution: Esri, HERE, De Lorme, MapmyIndia, (c) OpenStreetMap contributors, and the GIS user community. (<https://doc.arcgis.com/en/arcgis-online/reference/static-maps.htm>, accessed 22/02/2018)

Karl Sapper published an eruption catalogue in 1917. When he was satisfied there was sufficient evidence, Sapper categorised a number of historical eruptions by their products (lava and/or tephra) and their volume ($> 0.1 \text{ km}^3$ and $> 1 \text{ km}^3$). Sapper also cited Meyen's 1835 account but, reflecting Von Buch's interpretation, he took the existence of a 'great' (or 'powerful') eruption of 'Babuyan Claro' in 1831 as a fact (Sapper 1917; Table 1). On the basis of his reading of the Report of the Krakatoa Committee of the Royal Society (Symons 1888), however, he took a further step. Part IV, section V of that Report presents a list of observations of unusual atmospheric optical phenomena over the period 1500–1886, accompanied by a list of historical eruptions. In the case of 1831, alongside the unusual atmospheric optical phenomena seen in August, September and October 1831, six eruptions were listed including that of Babuyan Claro ("Babujan Islands (Great eruption)") (Symons 1888). The Report did not explicitly discuss the attribution of the phenomena to one eruption or another. Sapper, however, explicitly stated a

further assumption (A3) that the eruption of Babuyan Claro could have contributed to the unusual atmospheric optical phenomena observed in Europe (in August, September and October 1831) (Sapper 1917; Table 1) on the basis of which he assigned it a commensurate tephra volume of 0.1 to 1 km^3 (Sapper 1917, Table 1). This volume estimate has been widely used in subsequent quantitative treatments of the eruption.

For example, erupted volume (lava and/or tephra) is a key determinant in assigning a value for the Volcanic Explosivity Index (VEI): volumes of 0.01 to 0.1 km^3 are correlated with a VEI of 3 and 0.1 to 1 km^3 with a VEI of 4 (Newhall and Self 1982). Sapper's (1917) volumes were used to assign VEI values, although they were treated as 'rough estimates' and were not mechanically mapped: a number of early nineteenth century eruptions to which Sapper had assigned a 0.1 – 1 km^3 volume have been variously assigned VEI values of 3, 3? and 4 (Simkin et al. 1981; Newhall and Self 1982). The 1831 event was initially treated as an eruption of Smith volcano and assigned a VEI = 3?



Fig. 4 The volcanoes of Babuyan Claro island. Image reproduced with permission: Jonathan Torgovnik / Getty Images News / Getty Images. Smith volcano, in the foreground (known locally as 'Pokis', meaning 'bald' or alternatively as 'Mt. Babuyan'), underwent a series of eruptions between 1907 and 1924 (Maree 2005; GVP 2013). Babuyan Claro volcano ('Chinteb a Wasay', meaning 'cut of the axe' or alternatively 'Mt. Pangasun') is in the middle distance (Maree 2005; GVP 2013)

(Simkin et al. 1981). It was subsequently reclassified as an eruption of Babuyan Claro volcano (COMVOL 1981; Simkin and Siebert 1994) with a VEI of 4? (Simkin and Siebert 1994), bringing the value into line with Sapper's 0.1–1 km³ volume estimate and assigning to Babuyan Claro the largest VEI value of any eruption listed in 1831. It is noteworthy that the reclassification was contemporary with Zielinski et al.'s (1994) identification of this eruption as the source of the large magnitude Greenland sulphate peak. A VEI of 4? continues to be cited in the GVP database (GVP 2013, although see the accompanying note in Table 1).

Erupted volume also determines the Magnitude (M) of explosive eruptions, defined as (Pyle 2000):

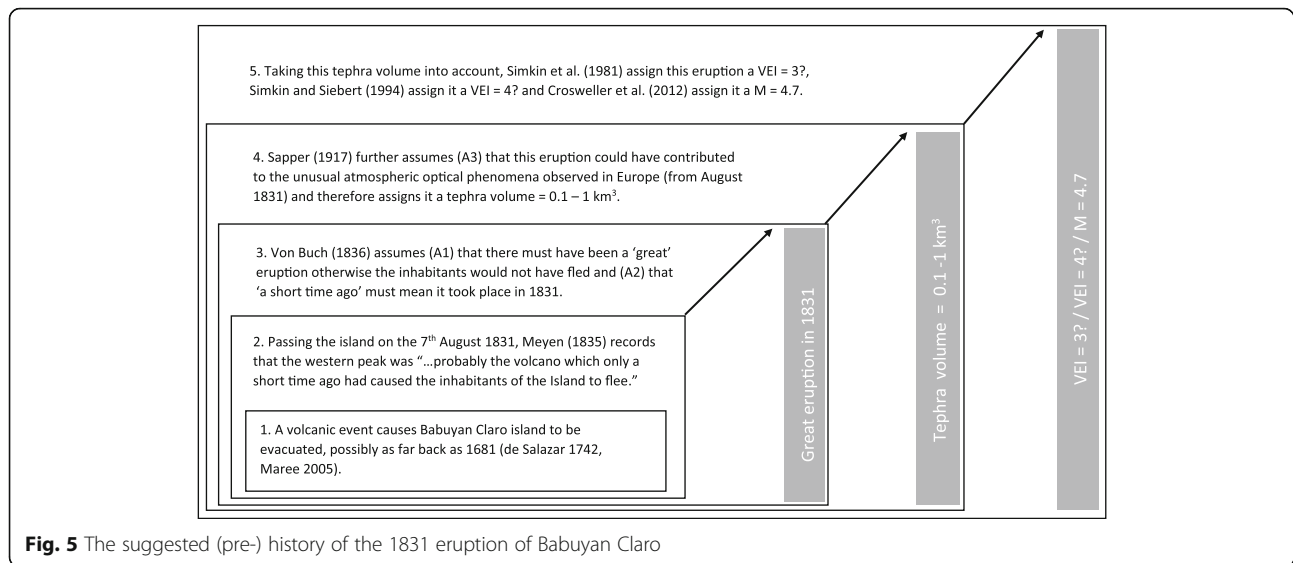
$$M = \log_{10}(\text{mass of erupted tephra or lava (kg)}) - 7$$

The LaMEVE database lists a Magnitude of 4.7 for this eruption, which is consistent with taking the mid-point of Sapper's 0.1–1 km³ volume and using a default tephra density of 1000 kg m⁻³ to obtain a mass of 5×10^{11} kg (Croweller et al. 2012). The VEI of 4? and Magnitude of

4.7 (Croweller et al. 2012; GVP 2013) are therefore essentially re-statements of Sapper's (1917) 0.1–1 km³ estimate for tephra volume.

The 1831 eruption of Babuyan Claro as a false eruption

Our analysis shows that there is no reason to believe that there was an 1831 eruption of Babuyan Claro. Although Meyen's (1835) account has long been treated as the single primary source for an 1831 eruption of Babuyan Claro, in fact it provides no evidence of any such eruption and, indeed, it is possible that it mistakenly refers to an event as far back as 1681. The records of such an eruption in the present-day literature can be explained as the result of Von Buch's (1836) and Sapper's (1917) mistaken assumptions about this account. These results are summarised in Fig. 5. We therefore suggest that the 1831 eruption of Babuyan Claro is a false event and that the records of it in the present-day literature are erroneous, including the VEI of 4? listed in the GVP database (GVP 2013) and the Magnitude of 4.7 listed in the LaMEVE database (Croweller et al. 2012).



Consequences for volcanology and climatology

Discounting the 1831 eruption, the remaining historical eruptions of volcanoes on Babuyan Claro island listed in the GVP database comprise, for Babuyan Claro volcano, 1860 (VEI = 2) and 1913 (VEI = 2, uncertain) and for Smith volcano, 1652 (VEI = 3) and 1907–1924 (five eruptions, all VEI = 2) (GVP 2013). The modest eruption reported by de Salazar (1742) in 1681 is not listed. In the absence of a historical eruption with a VEI \geq 4, Babuyan Claro might not be as violently explosive as previously thought: it may thus have a lower potential for triggering tsunamigenic hazards (Paris et al. 2014) and its inclusion in the LaMEVE database will no longer be warranted. Neither will there be any evidence to suggest that Babuyan Claro appears to form 'part of a high sulphur producing system' (Zielinski 2000).

The mis-identification of an eruption of Babuyan Claro as the source of the 1831 stratospheric sulphate aerosol will have hidden its actual source: at least one alternative eruption will have instead produced the Krakatoa-like unusual atmospheric optical phenomena observed around the world in August, September and October 1831 (Arago 1832; Kiessling 1888; Symons 1888), the surface temperature anomalies which occurred between 1831 and 1833 (White et al. 1997; Shindell et al. 2004; Fischer et al. 2007; McCarroll et al. 2013) and the large magnitude Greenland sulphate peak deposited during 1831–1833 (Zielinski et al. 1994; Gao et al. 2008; Sigl et al. 2013).

A small magnitude sulphate peak resulting from deposition during 1831–1833 has also been reported in ice-cores from Antarctica (Sigl et al. 2013). Although the detection of near synchronous (within c. 1 year) sulphate peaks in both Greenland and Antarctic ice-cores suggests a single low-latitude (tropical) eruption, as was thought to be the case with Babuyan Claro, another possibility is that

the sulphate peaks are the result of independent eruptions higher in the northern and southern hemispheres (Sigl et al. 2013). It has been suggested, for example, that the timing of the Antarctic sulphate peak instead likely points to a local Antarctic eruption in 1833 (Crowley and Unterman 2013). The alternative source of the Greenland sulphate peak could therefore be located at a mid- or high- latitude site in the northern hemisphere. If Babuyan Claro is discounted, the next largest magnitude eruptions listed for 1831 have a VEI of 3 (GVP 2013): Campi Flegrei Mar Sicilia, off the south coast of Sicily (37.1° N.) (also known as Ferdinanda, Giulia (Julia) and Graham Island); Guagua Pichincha, Ecuador (0.1° S.); and Mount St. Helens, U.S.A. (46.2° N.). A VEI of 3 is associated with only a 'possible' stratospheric injection of eruption products (Newhall and Self 1982) but this is more likely to happen at a mid- or high- latitude site than a low latitude site given the decrease in height of the tropopause from c. 15–17 km at the equator to c. 8–9 km at the poles: it has been estimated, for example, that the columns of some 23% of eruptions at mid- or high- latitude sites with a VEI of 3 may penetrate the tropopause (Pyle et al. 1996). Whether one of these known 1831 eruptions could be the source of the Greenland sulphate peak or whether an as yet unknown eruption is responsible is the subject of continuing work.

Although a detailed discussion is beyond the scope of this paper, a change of source eruption for the Greenland sulphate peak from one at a low-latitude site to one at a mid- or high- latitude site in the northern hemisphere could have quantitative consequences for the correct representation of the 1831 stratospheric sulphate aerosol in climate models. The mass (and therefore the forcing effect) of the 1831 stratospheric sulphate aerosol is typically reconstructed by 'scaling up' the magnitude of the

Greenland sulphate peak (Zielinski 1995; Gao et al. 2008; Arfeuille et al. 2014; Toohey and Sigl 2017). The choice of ‘scaling factor’ depends on the geographical location (latitude) of the assumed source eruption. Substituting a mid-latitude site in the northern hemisphere instead of the low-latitude site of Babuyan Claro, for example, would be expected to reduce the reconstructed mass by at least 25% (Zielinski 1995; Toohey and Sigl 2017).

The Babuyan Claro example is therefore distinct from other examples of misattributed eruptions where a climate impacting stratospheric sulphate aerosol has been produced, but where the change in source location involves only a small distance at similar latitude. For example, a VEI 5 eruption in 1641 (GVP 2013), which produced an even larger Greenland sulphate peak than that in 1831 (Zielinski et al. 1994; Sigl et al. 2013), had previously been attributed to Awu, in Indonesia, but Delfin et al. (1997) established that it had taken place at Parker volcano, also in the Philippines, only c. 300 km away.

Conclusions

The 1831 eruption of Babuyan Claro is regarded as one of the most significant volcanic climate forcing events of the nineteenth century (Zielinski 1995; Arfeuille et al. 2014; Toohey and Sigl 2017). However, our analysis shows that there is no reason to believe that there was an eruption of Babuyan Claro in 1831. The historical account which has long been treated as the single primary source for this eruption, Meyen (1835), in fact provides no evidence of any such eruption. The records of an 1831 eruption of Babuyan Claro in the present-day literature can be explained as the result of Von Buch’s (1835) and Sapper’s (1917) mistaken assumptions about this account. We therefore suggest that the 1831 eruption of Babuyan Claro is a false event and that its records in the present-day literature are erroneous, including the VEI of 4? listed in the GVP database (GVP 2013) and the Magnitude of 4.7 listed in the LaMEVE database (Crosweller et al. 2012).

The literature on false or misattributed eruptions is comparatively sparse. Two recent examples are Guidoboni (2010), discussing an example dating to 1198 which had been assigned a VEI of 1 and the above-discussed Delfin et al. (1997). The Babuyan Claro example therefore represents a significant further example of recent date and large magnitude. In the absence of a historical eruption with a $VEI \geq 4$, the hazard assessment of Babuyan Claro volcano will have to be re-evaluated and, for example, the inclusion of Babuyan Claro in the LaMEVE database will no longer be warranted.

At least one alternative eruption, potentially at a mid- or high- latitude site in the northern hemisphere and with a high sulphur yield, will have to be identified as the source of the 1831 stratospheric sulphate aerosol. Such a change of source eruption could have quantitative consequences

for the correct representation of the 1831 stratospheric sulphate aerosol in climate models. The Babuyan Claro example thus emphasises the potential for misattributed eruptions to have a significant impact on both volcanology and climatology and reinforces Guidoboni’s (2010) call for the careful analysis of primary historical sources in volcanology. Whether one of the remaining historical eruptions listed for 1831 in the GVP database could be the source or whether an as yet unknown eruption is responsible is the subject of continuing work.

Abbreviations

CAVW: Catalogue of the Active Volcanoes of the World; CDSB: Complete Dictionary of Scientific Biography; COMVOL: (Philippines) Commission on Volcanology; GVP: Global Volcanism Program; LaMEVE: Large Magnitude Explosive Volcanic Eruptions; PHIVOLCS: Philippine Institute of Volcanology and Seismology; VEI: Volcanic Explosivity Index

Acknowledgments

We would like to thank Montse Manzano Fernandez for providing the translation of de Salazar (1742), Ben Andrews (Smithsonian Institution) for information regarding the GVP database and Danielle Charlton (UCL Hazard Centre) for assistance with the preparation of Fig. 3. We would also like to thank Russell Blong, an anonymous second reviewer and Christopher Newhall for the improvements to the manuscript which resulted from their comments.

Funding

This work was self-funded (CG).

Availability of data and materials

All data generated or analysed during this study are included in this published article.

Authors’ contributions

CG carried out the analysis presented in this study and drafted the manuscript. CK and SE critically reviewed the manuscript and provided guidance and expertise on the volcanological context. All authors read and approved the final manuscript.

Competing interests

The authors declare that they have no competing interests.

Publisher’s Note

Springer Nature remains neutral with regard to jurisdictional claims in published maps and institutional affiliations.

Received: 27 February 2018 Accepted: 21 August 2018

Published online: 05 September 2018

References

- Alvir A (1928) Extrusives and Ejectamenta of some Philippine volcanoes. In: proc. III Pan Pac. Sci. Congr. 1926, Vol. 1, Tokyo, pp. 758–760.
- Arago F (1832) Notices scientifiques par M. Arago: des Comètes en général, et en particulier, de la Comète qui doit reparaitre en 1832 et dont la révolution Est de 6 ans 3/4 in: bureau des longitudes Annuaire pour l’An 1832. Bachelier Père et Fils, Paris, pp. 156–288. Available at: <http://gallica.bnf.fr>. (Gallica, Bibliothèque nationale de France).
- Arfeuille F, Weisenstein D, Mack H, Rozanov E, Peter T, Brönnimann S. Volcanic forcing for climate modeling: a new microphysics-based data set covering years 1600–present. *Clim Past*. 2014;10(1):359–75. <https://doi.org/10.5194/cp-10-359-2014>.
- Becker G. Report on the geology of the Philippine Islands. Washington: U.S. Geol. Survey 1899–1900, 21st Ann. Rep. Part 3; 1901.
- Buch V. Description physique des Îles Canaries, suivie d’une indication des principaux volcans du globe. Paris et Strasbourg: F.G. Levrault; 1836. Available at: <http://books.google.com>

- CDSB. Meyen, Franz Julius Ferdinand. In: Complete Dictionary of Scientific Biography, vol. 9. Detroit: Charles Scribner's Sons; 2008. p. 344–5. Available via Gale Virtual Reference Library Accessed 5 Jan. 2017.
- COMVOL. Catalogue of Philippine volcanoes and solfataric areas. Quezon City: Philippine Comm. Volc; 1981.
- Crowther HS, Arora B, Brown SK, Cottrell E, Deligne NI, Guerrero NO, Hobbs L, Kiyosugi K, Loughlin SC, Lowndes J, Nayemil M, Siebert L, Sparks RSJ, Takarada S, Venzke E. Global database on large magnitude explosive volcanic eruptions (LaMEVE). *J Appl Volcanol*. 2012;1(4):1–13. <https://doi.org/10.1186/2191-5040-1-4>.
- Crowley TJ, Unterman MB. Technical details concerning development of a 1200 yr proxy index for global volcanism. *Earth Syst Sci Data*. 2013;5(1):187–97. <https://doi.org/10.5194/essd-5-187-2013>.
- Daubeny C. A description of active and extinct volcanos, of earthquakes. In: And of thermal springs, 2nd edn. London: Richard and John E. Taylor; 1848.
- De Salazar V (1742) Historia de la Provincia de el Santissimo Rosario de Philipinas, China y Tunking, de el Sagrado orden de Predicadores: tercera parte. Manila. Available at: <http://bdh.bne.es>. (Biblioteca Digital Hispánica, Biblioteca Nacional de España).
- Defant MJ, Maury RC, Joron J, Feigenson MD, Leterrier J, Bellon H, Jacques D, Richard M. The geochemistry and tectonic setting of the northern section of the Luzon arc (the Philippines and Taiwan). *Tectonophysics*. 1990;183:187–205. [https://doi.org/10.1016/0040-1951\(90\)90416-6](https://doi.org/10.1016/0040-1951(90)90416-6).
- Delfin FG Jr, Newhall CG, Martinez ML, Salonga ND, Bayon FEB, Trimble D, Solidum R. Geological, ^{14}C and historical evidence for a 17th century eruption of Parker volcano, Mindanao. *Philippines J Geol Soc Philippines*. 1997;52(1):25–42.
- Fischer EM, Luterbacher J, Zorita E, Tett SFB, Casty C, Wanner H. European climate response to tropical volcanic eruptions over the last half millennium. *Geophys Res Lett*. 2007;34:L05707. <https://doi.org/10.1029/2006GL027992>.
- Fuchs C (1865) Die Vulkanischen Erscheinungen der Erde C. F. Winter'sche Verlagshandlung, Leipzig & Heidelberg. Available at: <http://books.google.com>.
- Fuchs C. *Vulcani e Terremoti*. Milan: Dumolard; 1881.
- Fuchs C. *Les Volcans et les Tremblements de Terre*, 4th edn. Paris: Félix Alcan; 1884.
- Gao C, Robock A, Ammann C. Volcanic forcing of climate over the past 1500 years: an improved ice-core based index for climate models. *J Geophys Res*. 2008;113:D23111. <https://doi.org/10.1029/2008JD010239>.
- Global Volcanism Program (GVP) (2013) *Volcanoes of the World*, v. 4.4.1. (Venzke, E (ed.)). Smithsonian Institution. <https://doi.org/10.5479/si.GVP.VOTW4-2013>. Accessed 22 Sept 2015.
- Guidoboni E. History and volcanology: dialogue overdue? The case of a false eruption in medieval Italy. *Eos, Trans Amer Geophys Union*. 2010;91(26):231. <https://doi.org/10.1029/2010EO260002>.
- Horsburgh J. *The India directory, or directions for sailing to and from the east indies*, 2nd edn. London; 1817.
- Von Humboldt A (1859) *Cosmos*. Gide, Paris.
- Von Humboldt A (1869) *Cosmos*. Harper Brothers, New York.
- Humphreys W. *Physics of the air*, 3rd edn. New York and London: McGraw Hill; 1940.
- Judd J. *Volcanoes: what they are and what they teach*. New York: D. Appleton; 1881.
- Kiessling J (1888) Untersuchungen über Dämmerungserscheinungen zur Erklärung der nach dem Krakatau-Ausbruch beobachteten atmosphärisch-optischen Störung. L. Voss, Hamburg. Available at: <http://www.e-rara.ch>. (ETH-Bibliothek Zürich).
- Können GP. The prodigious halo of the other Huygens. *Appl Opt*. 2015;54(4):B185–93. <https://doi.org/10.1364/AO.54.00B185>.
- Lamb H. Volcanic dust in the atmosphere; with a chronology and assessment of its meteorological significance. *Phil Trans R Soc A*. 1970;266:425–533. <https://doi.org/10.1098/rsta.1970.0010>.
- Landgrebe G. *Naturgeschichte der Vulcane*. Gotha: Justus Perthes; 1855.
- Longpré M-A, Stix J, Burkert C, Hansteen T, Kutterolf S. Sulfur budget and global climate impact of the a.D. 1835 eruption of Cosigüina volcano, Nicaragua. *Geophys Res Lett*. 2014;41:1–9. <https://doi.org/10.1002/2014GL061205>.
- Lynch DK, Livingston W. *Color and light in nature*, 2nd ed. Cambridge: Cambridge University Press; 2001.
- MacDonald GA. *Volcanoes*. New Jersey: Prentice-Hall Inc; 1972.
- Maree J. *The Ibatan: A Genealogy of the People of Babuyan Claro Island*. Philippines: Summer Institute of Linguistics; 2005.
- McCarroll D, Loader NJ, Jalkanen R, Gagen MH, Grudd H, Gunnarson BE, Kirchhefer AJ, Friedrich M, Linderholm HW, Lindholm M, Boettger T, Los S O, Remmele S, Kononov YM, Yamazaki YH, Young GHF, Zorita E. A 1200-year multiproxy record of tree growth and summer temperature at the northern pine forest limit of Europe. *The Holocene*. 2013;23(4):471–84. <https://doi.org/10.1177/0959683612467483>.
- Meinel A, Meinel M. *Sunsets, twilights and evening skies*. Cambridge: Cambridge University Press; 1983.
- Mercalli G. *I Vulcani attivi della terra*. Milan: Hoepli; 1907.
- Meyen FJF. *Phytomie*. Berlin; 1830.
- Meyen FJF. *Reise Um Die Erde*. Berlin; 1835. Available at: <http://books.google.com>
- Newhall C, Self S. The volcanic Explosivity index (VEI): an estimate of explosive magnitude for historical volcanism. *J Geophys Res Oceans*. 1982;87(C2):1231–8. <https://doi.org/10.1029/JC087IC02p01231>.
- Paris R, Switzer AD, Belousova M, Belousov A, Ontowirjo B, Whelley PL, Ulvrova M. Volcanic tsunami: a review of source mechanisms, past events and hazards in Southeast Asia (Indonesia, Philippines, Papua New Guinea). *Nat Hazards*. 2014;70(1):447–70. <https://doi.org/10.1007/s11069-013-0822-8>.
- Perrey A (1860) Documents Sur les tremblements de terre et les phénomènes volcaniques dans l'archipel des Philippines. Extrait des Mémoires de l'Academie de Dijon.
- PHIVOLCS (2008), Active Volcanoes. Philippine Institute of Volcanology and Seismology. <http://www.phivolcs.dost.gov.ph>. Accessed 12 Mar 2015.
- Pyle DM. In: Sigurdsson H, Houghton BF, SR MN, Rymer H, Stix J, editors. *Sizes of volcanic eruptions*. In *Encyclopedia of Volcanoes*. London: Academic Press; 2000.
- Pyle DM, Beattie PD, Bluth GJS. Sulphur emissions to the stratosphere from explosive volcanic eruptions. *Bull Volcanol*. 1996;57:663–71. <https://doi.org/10.1007/s004450050119>.
- Robock A. Volcanic eruptions and climate. *Rev Geophys*. 2000;38(2):191–219. <https://doi.org/10.1029/1998RG000054>.
- Saderra Masó M. Volcanoes and seismic centers of the Philippine Archipelago. Washington: Bulletin 3 (Census of the Philippine Islands), Department of Commerce and Labour (Bureau of the Census); 1904.
- Saderra Masó M. *Census of the Philippine Islands 1903*, vol. 1. Washington: United States Bureau of the Census; 1905.
- Saderra Masó M. *Active Philippine volcanoes*. Manila: Weather bureau, April; 1922.
- Saderra Masó M. *Active Philippine Volcanoes*. Manila: Seism. Bull. Philipp. Weather Bureau, Publ. 19; 1924.
- Saderra Masó M (1925) *Philippine Volcanoes*. *Verh. Geol. Mijnb. Gen. Ned. & Kol. Geol. Ser., Vol VIII*, Mouton & Co., Hague, pp. 379–386.
- Sapper K. *Katalog der geschichtlichen Vulkanausbrüche*. Straßburg: Trübner; 1917.
- Sapper K. *Vulkankunde*. Stuttgart: J. Engelhorn's Nachf; 1927.
- Schneider K. *Die Vulkanischen Erscheinungen der Erde*. Berlin: Gebrüder Borntraeger; 1911. Available at: <http://archive.org>
- Scrope G. *Volcanos...With a descriptive catalogue of all known volcanos and volcanic formations*, 2nd edn. London: Longman, Green, Longmans and Roberts; 1862.
- Semper C. *Die Philippinen und ihre Bewohner*, a. Würzburg: Stuber; 1869. Available at: <http://books.google.com>
- Shaw N. *Manual of meteorology*, 2nd edn. Cambridge: Cambridge University Press; 1936.
- Shindell DT, Schmidt GA, Mann ME, Faluvegi F. Dynamic winter climate response to large tropical volcanic eruptions since 1600. *J Geophys Res*. 2004;109:D05104. <https://doi.org/10.1029/2003JD004151>.
- Siebert L, Simkin T, Kimberly P. *Volcanoes of the world*, 3rd edn. Berkeley and Los Angeles: Univ. of California Press; 2010.
- Sigl M, McConnell JR, Layman L, Maselli O, McGwire K, Pasteris D, Dahl-Jensen D, Steffensen JP, Vinther B, Edwards R, Mulvaney R, Kipfstuhl S. A new bipolar ice core record of volcanism from WAIS divide and NEM and implications for climate forcing of the last 2000 years. *J Geophys Res Atmos*. 2013;118:1151–69. <https://doi.org/10.1029/2012JD018603>.
- Simkin T, Siebert L. *Volcanoes of the world*, 2nd edn. Tucson: Geoscience Press; 1994.
- Simkin T, Siebert L, McClelland L, Bridge D, Newhall C, Latter JH. *Volcanoes of the world*. Stroudsburg: Hutchinson Ross Publishing; 1981.
- Smith W. *Geology and mineral resources of the Philippine Islands*. Manila: Bureau of Printing; 1924.
- Symons GJ, editor. *The eruption of Krakatoa and subsequent phenomena*. London: Trübner & Co; 1888.
- Toohy M, Sigl M. Volcanic stratospheric sulfur injections and aerosol optical depth from 500 BCE to 1900 CE. *Earth Syst Sci Data*. 2017;9:809–31. <https://doi.org/10.5194/essd-9-809-2017>.
- Van Padang N. *Philippine Islands and Cochin China*. In: Part II, Catalogue of the Active Volcanoes of the World including Solfataric Fields. Naples: International Volcanological Association; 1953.
- Von Hoff K. *Chronik der Erdbeben und Vulcan-Ausbrüche*, Zweiter Theil. Gotha: Justus Perthes; 1841. Available at: <http://books.google.com>
- Von Humboldt A. *Kosmos*. Stuttgart and Tübingen: Gotha; 1858.

- White DE, White JW, Steig EJ, Barlow LK. Reconstructing annual and seasonal climatic responses from volcanic events since AD 1270 as recorded in the deuterium signal from the Greenland ice sheet project 2 ice core. *J Geophys Res Atmos*. 1997;102(D16):19683–94. <https://doi.org/10.1029/97DJ00774>.
- Zielinski G. Stratospheric loading and optical depth estimates of explosive volcanism over the last 2100 years derived from the Greenland ice sheet project 2 ice core. *J Geophys Res*. 1995;100(D10):20,937–55. <https://doi.org/10.1029/95JD01751>.
- Zielinski G. Use of paleo-records in determining variability within the volcanism-climate system. *Quat Sci Rev*. 2000;19:417–38. [https://doi.org/10.1016/S0277-3791\(99\)00073-6](https://doi.org/10.1016/S0277-3791(99)00073-6).
- Zielinski GA, Mayewski PA, Meeker LD, Whitlow S, Twickler MS, Morrison M, Meese DA, Gow AJ, Alley RB. Record of volcanism since 7000 BC from the GISP2 Greenland ice core and implications for the volcano-climate system. *Science*. 1994;264(5161):948–52. <https://doi.org/10.1126/science.264.5161.948>.

Submit your manuscript to a SpringerOpen[®] journal and benefit from:

- ▶ Convenient online submission
- ▶ Rigorous peer review
- ▶ Open access: articles freely available online
- ▶ High visibility within the field
- ▶ Retaining the copyright to your article

Submit your next manuscript at ▶ [springeropen.com](https://www.springeropen.com)
